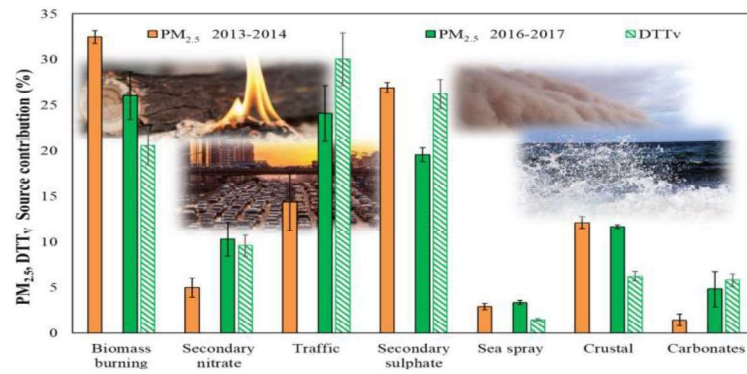


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## Graphical abstract



## Highlights

- Source contributions have inter-annual variability higher for  $PM_{2.5}$  than for  $PM_{10}$ .
- Inter-annual variability of crustal, carbonates, and sea spray were lower than 0.3% of PM.
- An increase of secondary nitrate and a decrease of sulphate in  $PM_{2.5}$  was observed.
- $DTT_V$  of  $PM_{2.5}$  peaked in cold periods, instead  $DTT_V$  in  $PM_{10-2.5}$  peaked in summer.
- $DTT_V$  in  $PM_{2.5}$  and in coarse ( $PM_{10-2.5}$ ) fractions were influenced by different sources.

# Inter-annual variability of source contributions to PM<sub>10</sub>, PM<sub>2.5</sub>, and oxidative potential in an urban background site in the central Mediterranean

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## Abstract

Airborne particulate matter (PM) is studied because of its effects on human health and climate change. PM long-term characterisation allows identifying trends and evaluating the outcomes of environmental protection policies. This work is aimed to study the inter-annual variability of PM<sub>2.5</sub> and PM<sub>10</sub> concentrations and chemical composition in an urban background site (Italy). A dataset of daily PM<sub>2.5</sub> and PM<sub>10</sub> was collected in the period 2016-2017, including the content of OC, EC, major water-soluble ions, main metals, and compared to a similar dataset collected in the period 2013-2014. Oxidative potential using DTT assay (dithiothreitol) was evaluated and expressed in DTT<sub>v</sub> as 0.39 nmol/min·m<sup>3</sup> in PM<sub>10</sub> and 0.29 in PM<sub>2.5</sub> nmol/min·m<sup>3</sup>. PM source apportionment was computed using the EPA PMF5.0 model and source contributions compared with those of a previous dataset collected between 2013 and 2014. Multi linear regression analysis identified which source contributed (p<0.05) to the oxidative potential of each size fraction. Inter-annual trends were more evident on PM<sub>2.5</sub> with reductions of biomass burning contribution and increases in traffic contribution in the 2016-2017 period. Crustal contributions were similar for the two periods, in both size fractions. Carbonates were comparable in PM<sub>10</sub> with a slight increase in PM<sub>2.5</sub>. Sea spray decreased in PM<sub>10</sub>. The DTT<sub>v</sub> of PM<sub>2.5</sub> peaked during cold periods, while, the DTT<sub>v</sub> of the PM<sub>10-2.5</sub> fraction peaked in summer, suggesting that different sources, with different seasonality, influence OP in the PM<sub>2.5</sub> and PM<sub>10-2.5</sub> fractions. Analysis showed that sea spray, crustal, and carbonates sources contribute ~13.6% to DTT<sub>v</sub> in PM<sub>2.5</sub> and ~62.4% to DTT<sub>v</sub> in PM<sub>10-2.5</sub>. Combustion sources (biomass burning and traffic) contribute to the majority of DTT<sub>v</sub> (50.6%) in PM<sub>2.5</sub> and contribute for ~26% to DTT<sub>v</sub> in PM<sub>10-2.5</sub>. Secondary nitrate contributes to DTT<sub>v</sub> in both fine and coarse fraction; secondary sulfate contribute to DTT<sub>v</sub> in PM<sub>2.5</sub> with negligible contributions to DTT<sub>v</sub> in PM<sub>10-2.5</sub>.

**Keywords:** oxidative potential, source apportionment, PMF, PM composition trend, DTT assay

## 1. Introduction

Aerosol pollution represents a concern for environments, ecosystems, and human health. It has been associated with millions of premature deaths, cardiovascular and respiratory morbidity worldwide each year (Shiraiwa et al., 2017; Ahmed et al., 2018; Hallquist, et al., 2009; Lelieveld et al., 2015; Pope et al., 2015; Cohen et al., 2017). Atmospheric pollution affects not only developing countries but also Europe, the United States, and Australia, which, despite the progress made in reducing industrial and transport emissions, are still facing this problem.

Human exposure to high PM<sub>10</sub> (where PM<sub>10</sub> is referring to particles with an aerodynamic diameter < 10 µm) levels has been associated to different adverse health effects involving the cardiovascular and respiratory systems, diseases like lung cancer, and various kind of allergies (Manigrasso et al., 2022 and references therein). Prompted by the previous observations, in the 1997 the US EPA promulgated the NAAQS (National Ambient Air Quality Standards) for PM<sub>2.5</sub>, adopting PM<sub>2.5</sub> as the indicator for fine particles (particles with an aerodynamic diameter < 2.5 µm) and identifying PM<sub>2.5-10</sub> as the coarse fraction of particles (i.e. particles having an aerodynamic diameter in the range 2.5-10 µm) (USEPA 2013). The coarse fraction, could have relevant health effects even it if is generally less studied compared to PM<sub>2.5</sub> (Adar et al., 2014). Roughly 4.2 million deaths were attributed to atmospheric pollution in 2016 by the World Health Organization (WHO). The majority of this mortality was due to exposure to fine particulate matter (PM<sub>2.5</sub>). Lelieveld et al. (2019) reported, for Europe, an annual excess mortality rate of 659,000 cases due to air pollution, calculated for the year 2015, again largely attributable to PM<sub>2.5</sub>. Furthermore, in Italy, the European Environment Agency (EEA, 2020) estimated that there were more than 50,000 premature deaths in 2018 attributable to long term exposure to PM<sub>2.5</sub>. The new air quality guidelines of the WHO reduced the annual limit of PM<sub>2.5</sub> from 10 µg/m<sup>3</sup> to 5 µg/m<sup>3</sup> (WHO, 2021). The toxicity of PM depends on several physical and chemical properties, most of them related to their sources, nevertheless, the effective mechanisms driving toxicity are not yet fully understood (Kelly et al., 2012; Kim et al., 2015; Rönkkö et al., 2021; Valavanidis et al., 2008; WHO, 2013). There are several evidences that some adverse health effects

78 could be driven by oxidative stress induced by exposure to PM (Ayres et al., 2008; Bates et al., 2019;  
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279 Guo et al., 2020; Li et al., 2015). As a consequence, several studies investigated the contributions of  
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580 sources to oxidative potential (OP) of PM, intended as an indirect metric of the ability of PM to induce  
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781 Reactive Oxygen Species (ROS) in biological systems (Costabile et al., 2019; Guascito et al., 2021;  
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1082 Massimi et al., 2020; Weber et al., 2021). A recent paper (Daellenbach et al., 2020) showed that the  
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1283 contributions of sources to PM mass concentrations and to OP are significantly different, suggesting  
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1584 that mitigation strategies aimed at reducing PM concentration alone could be not efficient in reducing  
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1785 oxidative potential. As a consequence, even if the link between OP and biological outcomes is still  
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1986 uncertain (Øvrevik, 2019; Lionetto et al., 2019; 2021; Jiang et al., 2019), it would be advisable to  
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2287 implement reduction strategies aimed to specific sources rather than to the total PM mass  
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2488 concentration and to investigate the long term trends of both PM concentration and composition.

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2689 In Europe, the emissions of a number of toxic atmospheric pollutants have been reduced during  
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2990 the last decades, with the aim to reduce the harmful effects of air pollution, in accordance with the  
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3291 2008 Directive on Ambient Air Quality and Cleaner Air for Europe and with the 2004 Directive on  
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3492 heavy metals and polycyclic aromatic hydrocarbons in ambient air. However, the assessment of the  
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3693 outcomes of these policies needs comprehensive studies of the impacts of natural and anthropogenic  
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3994 sources to the composition of the atmosphere. For this reason, long term measurements of particulate  
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4195 matter concentrations and chemical compositions are recommendable to obtain reliable assessments  
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4496 of the time trends allowing to estimate efficiency of the enforced mitigation strategies. Some studies  
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4697 have been done using both numerical models and long term ground level measurements in Europe  
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4998 (Waked et al., 2018; Yang et al., 2020) and in other areas of the World (Kyllönen et al., 2020; Jain et  
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5199 al., 2020; Friedman, 2020). Yang et al. (2020) used a global aerosol-climate model with an explicit  
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53100 source tagging approach (CAM5-EAST) to investigate source apportionment of PM in Europe  
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56101 between 1980 and 2018. They found a decrease of about 62% of total near surface sulphate, black  
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58102 carbon, and primary organic carbon due to the reductions of local emissions in Europe with an 8-9%  
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60103 contribution from the reductions in the Russia-Belarus-Ukraine area. In Italy, a slow reduction of

104 PM<sub>10</sub> and NO<sub>2</sub> levels were observed. This reduction is due to the joint decrease of primary particulate  
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105 emissions and of the main precursors of secondary particulate matter (i.e. nitrogen oxides, sulfur  
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106 oxides, ammonia) observed with the use of natural gas replacing fuels such as coal and oil, the  
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6  
107 introduction of catalysts in vehicles, and the adoption of technologies aimed at improving combustion  
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108 processes in industrial processes. However, exceedances of the PM<sub>10</sub> daily limit still occur in many  
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109 urban areas and in monitoring stations located near important road arteries (ISPRA, 2017).  
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14 In the scientific literature, many studies aimed to characterize PM properties “in situ” based on  
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171 PM measurements. However, most of these studies were based on sampling campaigns that lasted  
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192 from weeks to months, whereas long-term (several years or multi-decadal) measurements of PM  
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223 properties were done only in a limited number of sites (Rodríguez et al., 2012 and references therein).  
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244 Long term measurements are crucial for many reasons, such as, the assessment of the consistency of  
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26  
2715 remote sensing data (Chiapello et al., 1999; Smirnov et al., 2000); for model validations (Ginoux et  
28  
2916 al., 2004; Huneus et al., 2010); for identification of processes affecting the evolution of dust  
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327 emissions and transport (Prospero and Lamb, 2003).  
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348 The aims of this study are: (i) to investigate the inter-annual variability of the chemical  
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379 composition of PM<sub>2.5</sub> and PM<sub>10</sub>; (ii) using the positive matrix factorization model (PMF) to  
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3920 investigate the inter-annual trend of the contribution of the main natural and anthropogenic sources  
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4121 to the different PM fractions: PM<sub>2.5</sub>, PM<sub>10</sub> and the coarse PM<sub>10-2.5</sub>; (iii) to evaluate the contributions  
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4422 of the different sources to the oxidative potential of PM in three size ranges, estimated using the DTT  
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46  
4723 (Dithiothreitol) acellular assay. Particularly, only the PM water-soluble extract was analysed in term  
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4924 of OP (via DTT assay). This choice depends on observation that most of the literature OP data refers  
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5125 to the PM water-soluble fraction, that is considered more bioavailable (Gao et al., 2020a; Mukhtar et  
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5426 al., 2013). For example, many toxicological studies indicated the water-soluble metal as a possible  
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5627 harmful component of PM (Costa et al., 1997; Heal et al., 2005; Adamson et al., 2000; Ghio et al.,  
57  
58  
5928 2001), showing that water-soluble metals were able to generate reactive oxygen species (ROS), which  
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6129 switch on cellular pro-inflammatory response pathways in vitro and in vivo (Mukhtar et al., 2013).  
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130 About the application of the OP assays to insoluble PM, this is very difficult and it is still not clear  
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131 how much insoluble PM may contribute to the OP (Frezzini et al., 2022).  
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132 These aspects have been investigated using samples collected at an urban background Mediterranean  
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133 site: the Environmental-Climatic Observatory (ECO) of Lecce (South Italy). This observatory is a  
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134 regional station of the Global Atmosphere Watch program (GAW-WMO) and a facility of the  
10  
135 ACTRIS network, operative since 2013 (Dinoi et al., 2020; 2021a; 2021b; Cesari et al., 2018a;  
13  
136 2018b). Results provide data valuable for planning future mitigation strategies to reduce the impact  
15  
137 of atmospheric particulate matter on health in this area.  
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138

## 139 **2. Method**

### 140 **2.1. Measurement station and sampling**

141 PM<sub>10</sub> and PM<sub>2.5</sub> daily samples were collected at the Environmental-Climatic Observatory (ECO)  
27  
142 located inside the University Campus of Lecce (south Italy, 40°20'8''N– 18°07'28i''E, 37 m a.s.l.,  
30  
143 Figure S1). This Observatory is a regional station of the GAW- WMO network (Global Atmosphere  
32  
144 Watch) and has been already deeply described in other works (Cesari et al., 2018a, 2018b; Dinoi et  
35  
145 al., 2017).  
37

146 Between 15/11/2016 and 17/11/2017, daily samples were continuously collected at the  
40  
147 Observatory. Roughly, one sample every 3 days, for a total of 248 simultaneous samples (124 for  
42  
148 PM<sub>10</sub> and 124 for PM<sub>2.5</sub>), was selected and analysed for this work. The sampler used was a low-  
44  
149 volume (2.3 m<sup>3</sup>/h) dual channel sequential sampler (SWAM, Fai Instruments srl, Italy). The  
47  
150 instrument measures PM<sub>2.5</sub> and PM<sub>10</sub> concentrations using an automatic detection based on β-ray  
49  
151 attenuation with a typical uncertainty of 2% (PM<sub>10</sub>) and 3% (PM<sub>2.5</sub>) (Conte et al., 2020). Aerosol  
52  
152 mass was collected on quartz filters (Whatman Q-grade, diameter 47 mm) thermally pre-treated at  
54  
153 700 °C for 2 hours.  
57

### 154 155 **2.2. Chemical analysis**



156 The same approach of Cesari et al. (2018b) was used and it is here briefly described. Each filter was  
1  
157 divided into four quarters, three of them used for the chemical analysis and one used for replication  
3  
158 when needed.  
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159 A punch (1 cm<sup>2</sup>) of one quarter of filter was used for determination of OC and EC using the  
8  
160 thermo-optical transmittance (TOT) method (Sunset OC-EC Analyser), applying the EUSAAR2  
10  
161 protocol. The instruments was calibrated in the range 0-45 µgC/cm<sup>2</sup> using a sucrose standard solution.  
13  
162 The LODs (Limit Of Detections) were 0.03 µg/m<sup>3</sup> (EC) and 0.1 µg/m<sup>3</sup> (OC) and typical measurement  
15  
163 uncertainties were 5% for OC and 10% for EC (Merico et al., 2019). OC and EC concentrations were  
18  
164 determined subtracting contamination on blank filters.  
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21 The second quarter of filter was used to extract the water-soluble fraction to be devoted both to  
22  
23 oxidative potential analysis and concentrations of water soluble ions determination. The water-  
24  
25 soluble fraction of collected PM was extracted in 15 mL of deionized water (Milli-Q 18 M Ω) using  
26  
267 a 30-min sonication. A portion of each extract was used in High Performance Ion Chromatography  
28  
29  
30 a 30-min sonication. A portion of each extract was used in High Performance Ion Chromatography  
31  
32 (Dionex DX120 IC) system for determination of concentrations of major ions. The HPIC was  
33  
34  
35 equipped with an ED50 Conductivity detector, used with a 25 µL injection loop. Anions (Cl<sup>-</sup> NO<sub>3</sub><sup>-</sup>,  
36  
37 SO<sub>4</sub><sup>2-</sup>) were separated with a Dionex AS4A-4 column coupled with an IonPac AG14 guard column  
38  
39  
40 using as eluent 2.7 mM Na<sub>2</sub>CO<sub>3</sub> and 1.0 mM NaHCO<sub>3</sub> in isocratic mode. Cations (Na<sup>+</sup>, NH<sub>4</sub><sup>+</sup>, K<sup>+</sup>,  
41  
42 Mg<sup>2+</sup>, Ca<sup>2+</sup>) were separated with a Dionex CS12A-4 column with IonPac CG12A guard column,  
43  
44 using 20 mM MSA (methanesulfonic acid) as eluent in isocratic mode. The self-regenerating  
45  
46  
47 suppressors Dionex ASRS 300 for anions and Dionex CSRS 300 for cations were used. HPIC  
48  
49 calibration was done using single anions and cations solutions (Thermo Scientific\_Dionex\_IC  
50  
51  
52 Standard). The method detection limits (MDLs) values (g/L) were as follows: 8.5 (Na<sup>+</sup>), 7.0 (K<sup>+</sup>),  
53  
54 45.2 (NH<sub>4</sub><sup>+</sup>), 9.2 (Ca<sup>2+</sup>), 2.5 (Mg<sup>2+</sup>), 4.9 (Cl<sup>-</sup>), 10.5 (NO<sub>3</sub><sup>-</sup>), and 52 (SO<sub>4</sub><sup>2-</sup>).  
55

56  
57 Another portion of the water-soluble extracts was filtered with PTFE syringe filters (0.45 µm  
58  
59 porosity) to remove residual fibers and eventual insoluble materials. Successively, the filtered  
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61 aliquots were used to evaluate the oxidative potential (DTT assay) following the methodology of Cho  
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182 et al. (2005). Specifically, 3.5 mL of each extract was incubated at 37 ° C adding 0.5 mL of DTT  
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183 (1mM) and 1 mL of 0.5 M potassium phosphate buffer at pH 7.4 for times ranging from 5 to 90  
3  
184 minutes. At specific intervals (from 5 to 90 min), a 0.5 mL aliquot of the incubation mixture was  
4  
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185 collected and 0.5 mL of trichloroacetic acid (10% w/v) was added to stop the reaction. Afterwards, 2  
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9  
186 mL of 0.4M Tris-HCl, pH 8.9 containing 20 mM EDTA and 25 μL of 10 mM DTNB was added and,  
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11  
187 after 1 minute the absorbance of the solution was measured, at 412 nm, using a UV-Vis UVIKON  
13  
188 942 (KONTRON) spectrophotometer. The DTT depletion rate ( $\delta_{\text{DTT}}$ , pmol/min) was determined by  
15  
16  
189 linear regression between the measured absorbance and the time (Chirizzi et al., 2017). The detected  
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190 values of  $\delta_{\text{DTT}}$  were corrected using field blank measurements. The uncertainty on  $\delta_{\text{DTT}}$ , determined  
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291 by replication of measurements, was generally in the interval 3% - 15% (average ~ 6%). The DTT  
23  
292 depletion rate allows to calculate the OP in terms of  $\text{DTT}_V$ , i.e. activity normalized with the air volume  
25  
293 (V) actually sampled on each filter, or normalized with the mass of particulate matter collected on  
28  
294 the filter  $\text{DTT}_M$ .  
30

31  
195 The third quarter of filter was used in ICP-MS (NexIon 300X, Perkin Elmer, USA) to determine  
33  
396 concentration of different metals. Each sample was digested in closed Teflon vessels by a microwave  
35  
397 digester (Microwave Digestion System Start D, Milestone, Italy). The two steps procedure (EN  
37  
398 14385) was used. Firstly digestion in 1.2 mL HF (hydrofluoric acid, UltraTrace Analysis, 47%,  
40  
4199 Honeywell Fluka™) plus 1.8 mL HNO<sub>3</sub> (UltraTrace Analysis, 65-71%, Honeywell Fluka™) using  
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43  
200 temperature steps: from room temperature to 200 °C in 15 min, 15 min plateau at 200°C, and a final  
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46  
201 cooling back to room temperature. Successively, after addition of 14 mL of saturated H<sub>3</sub>BO<sub>3</sub> (Sigma  
47  
48  
202 Aldrich, ACS reagent, ≥99.5%), temperature was increased to 200 °C in 10 min, a 15 min plateau  
49  
50  
203 followed, and a final cooling to room temperature. Finally, samples were diluted to 25 mL in  
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53  
204 volumetric flask using deionized water (Milli-Q® 18.2 MΩ). Each sample (1 mL) was transferred  
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55  
205 into 10 mL volumetric flask, Rh (Fluka, Spectroscopic Grade) (final concentration 5 μg/L) as internal  
57  
58  
206 standard and Y (Fluka, Spectroscopic Grade) (final concentration 5 μg/L - Indium in case of rare  
59  
60  
61  
62  
63  
64  
65

207 earth elements analysis-) as second control were added and make up to the mark with 2% HNO<sub>3</sub>  
1  
208 solution. The ICP-MS was tuned using a Multi-Element Standard for Instrument Calibration solution  
3  
209 (Perkin Elmer) at 1 µg/L: Be, Ce, Fe, In, Li, Mg, Pb, U. The external calibration was performed for  
4  
5  
6  
210 the following elements: Al, As, Ba, Cd, Ce, Co, Cu, Dy, Fe, La, Li, Mn, Nd, Pb, Sb, Sr, Th, Ti, V,  
7  
8  
9  
211 Zn using rare earth elements mix (Perkin Elmer) and single analyte standard solutions (Fluka,  
10  
11  
212 Spectroscopic Grade). Quality control checks were performed monitoring the intensities of all  
12  
13  
213 internal standards for every sample analysis and analyzing the laboratory control samples (LCS), in  
14  
15  
214 each sample batch, at a frequency of one LCS every 10 samples. The LCS were pre-fired quartz filters  
16  
17  
215 digested using the same procedure used for samples and a mix of standards, selected to control both  
18  
19  
20  
216 matrix effect and possible memory effect. The MDLs, calculated with the same approach used for  
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Final concentrations of the different chemical species were corrected using the average level  
found in the blank samples. The concentration of a specific species in a samples was substituted with  
the threshold value  $\sigma_B/2$  (i.e. one half of the standard deviation in blank filters) when it was too low  
to be quantified. Uncertainties for measured concentrations were evaluated as described in Cesari et  
al. (2018b).

### 2.3. Source apportionment using Positive Matrix Factorization

The EPA PMF5 receptor model was used to identify the number of sources, their chemical  
profiles, and their contributions to the measured PM<sub>2.5</sub> and PM<sub>10</sub> concentrations. PMF is widely used  
worldwide for source apportionment being able to provide chemical profiles and contributions of  
sources, thereby operating with minimal information on the effective sources insisting on the site  
(Belis et al., 2020; Hopke et al., 2020). A single input data set was used, pooling together the chemical  
composition of PM<sub>2.5</sub> and PM<sub>10</sub> fractions to improve robustness of the results. The input variables  
were classified using both, the Signal-to-Noise (S/N) criteria (Paatero and Hopke, 2003) and the  
percentage of data above the detection limits (Amato et al., 2016; Contini et al., 2016). The species

233 OC, EC, Cl, SO<sub>4</sub><sup>2-</sup>, NH<sub>4</sub><sup>+</sup>, K<sup>+</sup>, Mg<sup>2+</sup>, Ca<sup>2+</sup>, Mn, Fe, As, and Ba, were classified as “strong” and used  
1  
234 directly in PMF5; while, the species NO<sub>3</sub><sup>-</sup>, Na<sup>+</sup>, Al, V, Co, Cu, Zn, Sr, Sb, Dy, La, Nd, and Th were  
3  
235 “weak”. The best solution for the base run was obtained using seven factors: biomass burning,  
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6  
236 ammonium nitrate, traffic, sulphate, sea spray, crustal, and carbonates. Compared to the previous  
8  
237 dataset, only the industrial source is missing likely because Pb (one of its main markers) was not  
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238 available in the “2016-2017” dataset. Successively, a constrained run was performed in order to  
11  
13  
239 improve the separation between factors profiles (Amato et al., 2016; Cesari et al., 2021). The  
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15  
240 constraints used were: pull down maximally OC in traffic factor; pull up maximally SO<sub>4</sub><sup>2-</sup> and Mg in  
16  
18  
241 crustal factor profile; pull down maximally OC in carbonate factor; pull up maximally NH<sub>4</sub><sup>+</sup> in  
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242 secondary sulphate factor. The final dQ change, compared to the base run, was 7.8%. Uncertainty  
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243 estimates in PMF results were obtained with the bootstrap method (Paatero et al., 2014). The  
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25  
244 bootstrap of the “constrained solution” (applied with 100 runs with random seed, block size suggested  
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27  
245 23, and R=0.6) gave a good mapping of the solution with unmapped cases limited to 1% and 2% for  
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246 traffic and nitrate (respectively) and to 6% (crustal) and 4% (carbonates).  
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33  
247 Successively, a multi-linear regression (MLR) analysis was done between the daily  
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35  
248 contributions of the different sources estimated by the model PMF and the daily measured DTT<sub>v</sub>  
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37  
249 values. This allowed to determine the contribution of each source identified by the PMF to the  
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250 measured PM oxidative potential. The MLR was done using the XLSTAT tool imposing the intercept  
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251 equal to zero and maintaining sources with statistical significance at 5% level (i.e. p<0.05). The  
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44  
45  
252 approach is the same already used in a previous work (Cesari et al., 2019).  
46  
47

48  
253 Finally, the differences observed between the seasonal periods considered (2013-2014 and  
49  
50  
254 2016-2017) in terms of PM source contributions estimated by PMF and by stoichiometric calculations  
51  
52  
255 were performed using the Analysis Tool Pack of Microsoft Excel. The ANOVA (Analysis of  
53  
54  
256 Variance) test was used to test the significance of the differences observed and the threshold of the  
55  
56  
57  
257 p-value was set at 5% (p-value 0.05) for the statistical test. All the significative differences founded  
58  
59  
258 with this test were shown with an asterisk in the figures reported in the results.  
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### 3. Results and discussion

#### 3.1. Inter-annual variabilities of the different chemical species

Average ( $\pm$ standard deviations) concentrations of PM, elements, water soluble ions, and carbonaceous species detected in the period 2016-2017 are reported in Table S1 and compared with the results found in the previous dataset (collected in the period 2013-2014). The yearly average concentration was 18.0 ( $\pm$ 11.4 standard deviation)  $\mu\text{g}/\text{m}^3$  for the fine fraction ( $\text{PM}_{2.5}$ ) and 26.9 ( $\pm$ 12.7 standard deviation)  $\mu\text{g}/\text{m}^3$  for the  $\text{PM}_{10}$  fraction. These values were comparable with the previous ones. The different measured species represent, on average, similar fractions of PM in the two periods: 55% of  $\text{PM}_{10}$  in 2016-2017 and 54% in 2013-2014; 60% of  $\text{PM}_{2.5}$  in both measurement periods, 2016-2017 and 2013-2014.

The main components of PM are compared in Table S1 showing that the sums of metals were comparable in the two periods representing roughly 2-2.5% of PM in both size fractions. The sums of ions were slightly lower in the new dataset: 28.6% in 2016-2017 and 29.5% in 2013-2014 for  $\text{PM}_{10}$ ; 26.7% in 2016-2017 and 27.8% in 2013-2014 for  $\text{PM}_{2.5}$ . Total carbon  $\text{TC}=\text{OC}+\text{EC}$  was comparable in the two measurement periods (for both size fractions), however, in 2016-2017 there was a slight shift with lower OC and larger EC compared to the previous period, even if considering the standard deviations, these differences were not statistically significant.

The analysis of the trends of the different chemical species (considering their averages and standard errors for the two period considered) showed comparable values of  $\text{K}^+$  ion, often associated to biomass burning emissions. The same happened for the ions  $\text{Na}^+$ ,  $\text{Cl}^-$  and  $\text{Mg}^{2+}$ , commonly considered of marine origin that could be a relevant source in this area (Contini et al., 2010). The slightly lower  $\text{Cl}^-$  observed in  $\text{PM}_{10}$  for the first period could be due to a different chlorine depletion in sea spray that is a relevant process observed at this site (Contini et al., 2014). The species associated to the secondary inorganic aerosol,  $\text{NH}_4^+$ ,  $\text{SO}_4^{2-}$  and  $\text{NO}_3^-$ , showed different trends in the two

285 fractions:  $\text{NH}_4^+$  showed a strong reduction in the second period of measurements (2016-2017) in both  
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286 fractions, contrarily,  $\text{SO}_4^{2-}$  showed a limited reduction. Finally,  $\text{NO}_3^-$  showed an increase in both size  
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287 fractions during the second measurement period. This could be interpreted with an increase of  
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288 ammonium nitrate limited to  $\text{PM}_{2.5}$ , because nitrate in  $\text{PM}_{10}$ , at this site, is mainly due to sodium  
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289 nitrate (Cesari et al., 2018b). The small reduction in secondary sulphate could be driven by the more  
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290 limited availability of ammonium. Elements of crustal origin, particularly Al and Fe showed no  
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291 trends, having similar concentrations during the two periods in both size fractions.  $\text{Ca}^{2+}$ , that at this  
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292 site is mainly related to carbonates due to re-suspended soil (Contini et al., 2014), was comparable in  
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293  $\text{PM}_{10}$  but presented an increase in  $\text{PM}_{2.5}$  during the 2016-2017 period. V, Co, and As had larger  
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294 concentrations in the 2016-2017 period and also larger standard deviations, however, the median  
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295 values were comparable in the two measurement periods. Dy showed larger concentrations in the  
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296 2016-2017 period. All the other elements detected in both fractions showed comparable  
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297 concentrations in the two sampling periods.

298 Mass closure approach and OC/EC ratio have been used to evaluate yearly variabilities of  
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299 secondary organic and inorganic aerosol, marine and crustal contributions to PM mass collected. The  
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300 OC/EC minimum ratio method was used to estimate the secondary organic carbon (SOC)  
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301 concentrations applying the equation reported in Pio et al., (2011), based on the  $(\text{OC}/\text{EC})_{\text{min}}$ . This  
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302 average ratio was  $5.9 (\pm 2.4)$  in  $\text{PM}_{10}$  and  $5.2 (\pm 2.0)$  in  $\text{PM}_{2.5}$ , in reasonable agreement with the ratios  
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303 observed in the first period of measurements (2013-2014) within the uncertainty intervals, that were  
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304  $7.8 (\pm 3.9)$  in  $\text{PM}_{10}$  and  $8.8 (\pm 4.9)$  in  $\text{PM}_{2.5}$  (Cesari et al., 2018b). As stated in Cesari et al. (2018b,  
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305 and references therein), the large OC/EC values observed were compatible with an urban background  
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306 site mainly influenced by biomass burning. The  $(\text{OC}/\text{EC})_{\text{min}}$  ratios were determined separately for  
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307  $\text{PM}_{10}$  and  $\text{PM}_{2.5}$ , looking at the minimum slopes, that was equal to 2 for both fractions (Figure S2).  
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308 By calculations, SOC was  $3.28 \mu\text{g}/\text{m}^3$  representing 62% of total OC in  $\text{PM}_{10}$ , and  $2.70 \mu\text{g}/\text{m}^3$   
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309 representing 56% of total OC in  $\text{PM}_{2.5}$ . The non-sea-salt sulphate was evaluated as  $\text{nss-SO}_4^{2-} = \text{SO}_4^{2-}$   
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610  $- 0.25\text{Na}^+$ , and represented 91% of total sulphate in  $\text{PM}_{10}$  and 97% in  $\text{PM}_{2.5}$ . The correlation between  
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311  $\text{SO}_4^{2-}$  and  $\text{NH}_4^+$ , 0.55 ( $p < 0.05$ ) and 0.61 ( $p < 0.05$ ) for  $\text{PM}_{10}$  and  $\text{PM}_{2.5}$ , respectively, indicated likely  
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312 the presence of a certain amount of ammonium sulphate in the sampled PM, estimated as  $2.6 (\pm 1.6)$   
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313  $\mu\text{g}/\text{m}^3$  in  $\text{PM}_{10}$  and  $2.5 (\pm 1.4) \mu\text{g}/\text{m}^3$  in  $\text{PM}_{2.5}$ . The correlations between ion  $\text{Na}^+$  with  $\text{Cl}^-$  and with  
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314  $\text{Mg}^{2+}$ , respectively 0.61 and 0.66, typically indicates the presence of sea spray (Querol et al., 2001).  
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315 This contribution was evaluated as  $\text{Cl}^- + 1.4468 * \text{Na}^+$  (Marenco et al., 2006) and was  $2.1 (\pm 1.8)$   
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316  $\mu\text{g}/\text{m}^3$  for  $\text{PM}_{10}$  and  $0.6 (\pm 0.4) \mu\text{g}/\text{m}^3$  for  $\text{PM}_{2.5}$ , with a major concentration in  $\text{PM}_{10}$  fraction as  
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317 expected. The crustal contribution was calculated using the formula  $1.15(1.89\text{Al} + 2.14\text{Si} + 1.67\text{Ti} +$   
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318  $1.4\text{Ca}^* + 1.2\text{K}^* + 1.36 \text{Fe})$  that consider metal oxides of Al, Si and Fe, plus the insoluble fraction of  
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319 K and Ca indicated with an asterisk (Cesari et al., 2018b). Carbonates were calculated from non-sea-  
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320 salt calcium and magnesium as  $1.5 \text{nss-Ca}^{2+} + 2.5 \text{nss-Mg}^{2+}$  (Perrino et al., 2014). The crustal  
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321 contribution, together with the carbonates, were  $3.25 (\pm 3.32) \mu\text{g}/\text{m}^3$  and  $2.14 (\pm 1.84) \mu\text{g}/\text{m}^3$   
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322 respectively for  $\text{PM}_{10}$ , while in  $\text{PM}_{2.5}$  were  $1.87 (\pm 2.15) \mu\text{g}/\text{m}^3$  and  $1.06 (\pm 1.19) \mu\text{g}/\text{m}^3$ .  
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303 Figure 1 shows the average concentrations of secondary organic carbon (SOC), secondary  
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324 inorganic aerosol (SIA, sum of nitrate, ammonium, and sulphate), sea spray, crustal, and carbonates  
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325 in absolute and relative terms, for the two sampling periods and for the two fractions  $\text{PM}_{10}$  and  $\text{PM}_{2.5}$ .  
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326 Considering the  $\text{PM}_{10}$  fraction, while contributions of natural origin, sea spray, crustal and carbonate  
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327 were similar, within the uncertainty ranges, a difference, even if not significant, was observed for  
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328 secondary organic and inorganic aerosol contributions, with a decreasing trend for both of them.  
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329 These differences could be mainly due to different atmospheric conditions (such as temperature,  
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330 relative humidity, presence of precursors) that could lead to different concentrations of secondary  
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331 aerosols in air. Looking at the  $\text{PM}_{2.5}$  fraction, marine and crustal sources provided similar  
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332 concentrations in the two period and a decrease of SOC and SIA was observed as in  $\text{PM}_{10}$  but is  
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333 statistical significant only for SOC. Carbonates showed a significant difference, with an  
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334 increasing trend compatible with the large  $\text{Ca}^{2+}$  concentrations observed in 2016-2017 limited to  
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335  $\text{PM}_{2.5}$ .  
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### 3.2. Trends of the contributions of sources

The best PMF5 solution was obtained using seven factors that were similar to those previously identified in Cesari et al. (2018b). The source profiles were shown in Figure 2. PMF5 reconstructed reasonably well measured concentrations for both size fractions (Fig. 2h). The portion unexplained by the model (i.e. the average difference between measured and reconstructed concentrations) was +6% for PM<sub>10</sub> and -6% for PM<sub>2.5</sub>. The linear fit of the correlation between measured and reconstructed PM concentrations had a slope of 1.06 ( $\pm 0.02$ ), an intercept of -1.1 ( $\pm 0.6$ )  $\mu\text{g}/\text{m}^3$  and a determination coefficient  $R^2$  of 0.89.

Figure 3 reports the estimated contributions, in absolute and relative terms, for the factors/sources identified and a comparison with the previous results obtained for the 2013-2014 period. In addition, Fig. 3 includes the sources evaluated with the calculations based on measured concentrations (Section 4.1 and indicated as “mass closure” values). As showed in Fig. 3, the differences observed are statistically significative only for PM<sub>2.5</sub> fraction (i.e. secondary nitrate, traffic, secondary sulphate and carbonates). There was a small reduction of biomass burning contribution, more evident on PM<sub>2.5</sub> with respect to PM<sub>10</sub> and an increase of the traffic contribution, again more relevant for PM<sub>2.5</sub>. It is noteworthy that the sum of these two main combustion sources was essentially constant in the two measurement periods and for the two size fractions. This could indicate a difficulty in separating the contributions of these two sources in the PMF source apportionment approach (Cesari et al., 2021). Crustal contributions were similar for the two periods and in the two size fractions, while, carbonates were comparable in PM<sub>10</sub> but a slight increase was observed in PM<sub>2.5</sub> in agreement with the increase of concentrations of Ca<sup>2+</sup> observed in this size fraction (Section 4.1). Sea spray decreased in PM<sub>10</sub> but not in PM<sub>2.5</sub>. This difference was probably due to the variability of atmospheric conditions, i.e. wind intensity and direction that could affect strongly sea spray generation and transport at this site (Contini et al., 2010). Secondary nitrate in the two measurement periods was comparable in PM<sub>10</sub> but an increase was observed in the fine fraction



362 (PM<sub>2.5</sub>) together with a decrease of secondary sulphate. This is compatible with a partial shift from  
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363 ammonium sulphate to ammonium nitrate in 2016-2017, as it was hypothesized in Section 4. This  
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364 shift involving ammonium is more relevant in PM<sub>2.5</sub> because a previous study at this site (Cesari et  
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365 al., 2018b) showed that nitrate in fine fraction is mainly ammonium nitrate especially in the cold  
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366 periods, while sodium nitrate is the most relevant form of nitrate for the coarse fraction (PM<sub>10-2.5</sub>). It  
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367 is interesting to observe that SIA (sum of secondary nitrate and sulphate) was essentially constant in  
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368 the two periods and in the two size fractions, within the variabilities indicated in Fig. 3.  
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369 The comparison of PMF contributions with the calculations based on measured concentrations  
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370 (mass closure on Fig. 3) were in reasonable agreement for sea spray, crustal, and carbonates; however,  
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371 there was an overestimation of SIA using PMF. This suggests that the chemical profiles of secondary  
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372 sources obtained by PMF were not purely stoichiometric nitrate and sulphate. This difference was  
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373 observed also in other sites in central Italy (Cesari et al., 2016) and it can be particularly relevant for  
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374 the contribution of secondary nitrate.  
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### 376 3.3. Influence of chemical composition and sources on oxidative potential

377 In Table S2, the yearly average of DTT<sub>V</sub> and DTT<sub>M</sub> detected in PM<sub>10</sub> and PM<sub>2.5</sub> were compared with  
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378 values measured in other urban background sampling sites of Italy and Europe. The values observed  
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379 at the site studied were comparable, within the variability range, with the levels observed in the other  
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380 sites.  
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381 Figure 4a shows the yearly and seasonal averages of DTT<sub>V</sub> in the PM<sub>2.5</sub> and in the coarse  
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382 fraction. The lower values of DTT<sub>V</sub> observed in the coarse fraction, compared to that of fine PM  
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383 fraction, were compatible with results discussed in other studies (Massimi et al., 2020;  
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384 Paraskevopoulou, et al. 2019; Simonetti et al., 2018). DTT<sub>V</sub> of PM<sub>2.5</sub> peaked during the cold period,  
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385 likely as a consequence of higher combustion emissions and of shallower boundary layer height that  
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386 concentrates ground-level aerosol, suggesting a possible association of OP with local winter sources  
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387 (such as biomass burning and traffic). DTT<sub>V</sub> in the PM<sub>10-2.5</sub> fraction followed a different seasonal  
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388 trend with the largest average value observed in summer suggesting that different sources influenced  
389 OP in the two size fractions. This was confirmed by an analysis of Pearson correlation coefficients  
390 because  $DTT_V$  in  $PM_{2.5}$  had relevant ( $p < 0.05$ ) correlations with the species OC (Pearson 0.69), EC  
391 (0.69),  $NO_3^-$  (0.46), and  $K^+$  (0.62) that are commonly associated to combustions sources emissions  
392 (Matawle et al., 2015; Callén et al., 2013). While, negligible correlations were observed with soil and  
393 crustal elements.  $DTT_V$  in the coarse fraction had lower correlation coefficients, compared to  $PM_{2.5}$ ,  
394 with OC and EC ( $\sim 0.3$ ) and larger coefficients, again in comparison with  $PM_{2.5}$ , for crustal and soil  
395 elements Al, Fe,  $Ca^{2+}$  (between 0.2 and 0.4).

396 Figure 4b shows the yearly and seasonal values of the specific oxidative potential (i.e.  $DTT_M$ )  
397 for  $PM_{2.5}$ ,  $PM_{10}$ , and  $PM_{10-2.5}$ . The  $DTT_M$  had a seasonal trend similar to that of  $DTT_V$  peaking during  
398 the cold period (winter and autumn). The  $DTT_M$  of the coarse fraction was quite relevant, being about  
399 70% (on average) of that calculated for  $PM_{2.5}$ , even if the  $DTT_V$  was about 2.9-time lower than that  
400 observed on  $PM_{2.5}$ . This was compatible with the results obtained in other sites (Hu et al., 2008).  
401 Considering studies on PM oxidative DTT potential trends reported in literature, are evident the high  
402 differences in OP trends detected in different sampling sites around the world. For example, in  
403 Nanjing (Yang et al., 2021) no obvious seasonal difference of  $OP_V$  was observed for outdoor  $PM_{2.5}$ ,  
404 while the outdoor  $OP_M$  exhibited a significant seasonal variation, with higher values in autumns and  
405 summer and the lowest in winter (Yang et al., 2021). In Hong Kong (Cheng et al., 2021) the OP trend  
406 was clearly correlated to meteorological conditions, being higher on regional days than on LRT and  
407 local days. Where the sampling days were classified into three categories mainly based the air mass  
408 backward trajectories, which are regional days (air quality in Hong Kong mainly influenced by  
409 continental air masses from the Pearl River Delta (PRD) region), long-range transport (LRT) days  
410 (mainly affected by northeastern coastal air masses), and local days (mainly under influence of local  
411 emission sources). Conversely, in another study performed in Atlanta (Gao et al., 2020b) there was  
412 no significant seasonal variation observed for either volume or mass- normalized OP water soluble  
413 DTT. In contrast, both volume and mass-normalized  $OP_{total}$ - DTT and thus OP water insoluble- DTT

414 varied between seasons, with higher levels in the cold months (Jan–Mar, Dec) than in summer (Jun–  
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415 Aug).

416 The contributions to OP of the PM sources identified were estimated using the MLR approach,  
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417 assuming that the DTT<sub>v</sub> was a linear combination of the PM mass concentrations due to the different  
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418 sources obtained with PMF5 (Weber et al., 2018). The slopes (i.e. fitting coefficients) represent the  
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419 intrinsic contributions of each source (expressed as nmol/min·µg). MLR results had  $R^2 = 0.88$  ( $R^2$   
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420 correct = 0.87) and a RMSE (root mean standard error) equal to 0.14. The model gave a reasonable  
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421 fit of the variables considered, as shown in Fig. 5(c). Table S3 reports the results of MLR analysis,  
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422 in terms of intrinsic contributions, standard error, p-values at 95% confidence intervals showing, for  
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423 each source investigated, that the MLR model gave a good fit for all sources.  
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424 Figure 5 shows the contributions, in absolute and in relative terms, of the different sources to  
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425 DTT<sub>v</sub> activity. MLR results showed that the different sources contribute differently to DTT<sub>v</sub> in fine  
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426 and coarse fraction confirming the hypothesis obtained by the analysis of the seasonal trends and  
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427 correlations. Sea spray, crustal, and carbonates, that are mainly sources of natural and soil origin at  
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428 this site (Cesari et al., 2018; Contini et al., 2014) had a small relative contribution (~13.6% in total)  
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429 to DTT<sub>v</sub> in PM<sub>2.5</sub> but they represented the majority of the contribution (~62.4% in total) of DTT<sub>v</sub> in  
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430 the coarse fraction. Contrarily, the main combustion sources (biomass burning and traffic) contributed  
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431 to the majority of DTT<sub>v</sub> (about 50.6%) in PM<sub>2.5</sub> but contributed for only ~26% to DTT<sub>v</sub> in the coarse  
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432 fraction. The low contribution of natural sources to DTT<sub>v</sub> compared to combustion sources in PM<sub>2.5</sub>  
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433 or in PM<sub>10</sub> was compatible with the results of other studies (Liu et al., 2018; Cesari et al., 2019;  
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434 Paraskevopoulou et al., 2019). Secondary nitrate contributed to DTT<sub>v</sub> in both fine and coarse  
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435 fractions, while secondary sulphate contributed to DTT<sub>v</sub> in PM<sub>2.5</sub> but had a negligible contribution  
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436 to DTT<sub>v</sub> in the PM<sub>10-2.5</sub> fraction. The contribution of nitrate was observed also in other studies  
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437 (Cesari et al., 2019; Paraskevopoulou et al., 2019) and it is likely due to the combustion origin of  
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438 nitrate precursors, for example from traffic and biomass burning. The MLR analysis showed also  
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439 high contributions to DTT<sub>v</sub> from secondary sulphate. Ammonium sulphate is not considered as a

440 strong redox agent and in some works (Cesari et al., 2019; Fang et al., 2016) secondary sulphate gave  
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441 an insignificant contribution to DTT<sub>v</sub>. However, in other studies a relevant contribution of secondary  
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442 sulphate to DTT<sub>v</sub> was observed (Liu et al., 2018; Massimi et al., 2020; Paraskevopoulou et al., 2019).  
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443 Results of this work showed a contribution from secondary sulphate to DTT<sub>v</sub> in PM<sub>2.5</sub> but a negligible  
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444 contribution to DTT<sub>v</sub> in PM<sub>10-2.5</sub>. This could be due to the presence of a certain amount of secondary  
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445 organic sulphate, in this source profile, that could influence and increase the measured DTT activity  
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446 of this source (Hakimzadeh et al., 2020; Verma et al., 2015).  
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#### 447 18 448 **4. Conclusions**

449 This research aimed to connect specific sources of PM to changes in seasons, as well as possible  
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450 adverse health outcomes. The inter-annual trends of the contributions of the main sources to PM<sub>2.5</sub>  
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451 and PM<sub>10</sub> and to their oxidative potential (in terms of DTT<sub>v</sub>) were assessed at the Environmental-  
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452 Climate Observatory (ECO) of Lecce (South Italy).  
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453 Results showed comparable values of total carbon for both PM<sub>2.5</sub> and PM<sub>10</sub>, with a decrease  
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454 of OC and an increase of EC. Comparable values in the two periods were observed for ions K<sup>+</sup>, Na<sup>+</sup>,  
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455 Cl<sup>-</sup> and Mg<sup>2+</sup>. Between 2013 and 2017, an increase of ammonium nitrate and a small decrease of  
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456 ammonium sulphate in PM<sub>2.5</sub> was observed. However, the total SIA was essentially constant in both  
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457 PM<sub>2.5</sub> and PM<sub>10</sub>. Elements of crustal origin Al, Fe showed no trends, but there was an increase of  
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458 Ca<sup>2+</sup> only in PM<sub>2.5</sub>.  
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459 Results of source apportionment showed that from 2013 to 2017 there is a reduction of the  
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460 biomass burning contribution and an increase of traffic contribution. This was more relevant for PM<sub>2.5</sub>  
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461 compared to PM<sub>10</sub>. However, the sums of these two main combustion sources were comparable in  
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462 the two periods. This could indicate a difficulty in separating the contributions of these two sources  
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463 in the application of PMF5. Crustal contributions were similar for the two periods in both PM<sub>2.5</sub> and  
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464 PM<sub>10</sub>, while carbonates were comparable in PM<sub>10</sub> but a slight increase was observed in PM<sub>2.5</sub>. Sea  
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465 spray decreased in PM<sub>10</sub>, likely because of the variability of atmospheric conditions.  
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466 The DTT<sub>v</sub> of PM<sub>2.5</sub> peaked during the cold period, as a consequence of higher contributions  
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467 of combustion sources. The DTT<sub>v</sub> in the PM<sub>10-2.5</sub> fraction followed a different seasonal trend with  
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468 maximum average value observed in summer, suggesting that different sources, having different  
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469 seasonality, influenced OP in these two size fractions. The DTT<sub>v</sub> of PM<sub>10-2.5</sub> was, on average,  
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470 significantly lower than that of PM<sub>2.5</sub>. The MLR analysis of sources contributing to DTT<sub>v</sub> showed  
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471 that natural sources (i.e. sea spray, crustal and carbonates) have a small relative contribution (~13.6%  
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472 in total) to DTT<sub>v</sub> in PM<sub>2.5</sub> but they represented the major contributors (~62.4% in total) to DTT<sub>v</sub> in  
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473 the PM<sub>10-2.5</sub> fraction. Conversely, combustion sources (biomass burning and traffic) contributed to  
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474 the majority of DTT<sub>v</sub> (about 50.6%) in PM<sub>2.5</sub> but contribute for only ~26% to DTT<sub>v</sub> in the PM<sub>10-2.5</sub>  
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475 fraction. Secondary nitrate contributed to DTT<sub>v</sub> in both size fractions, while secondary sulphate  
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476 contributed to DTT<sub>v</sub> in PM<sub>2.5</sub> with a negligible contribution to DTT<sub>v</sub> in the PM<sub>10-2.5</sub> fraction.  
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477 The results could be extended in future studies to investigate trends on longer periods and  
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478 could provide valuable information for planning future mitigation strategies aimed to reduce the  
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479 impact of atmospheric particulate matter on health in this area.  
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492 **Declaration of interests**

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494 The authors declare that they have no known competing interests.

495  
496 **Author contributions.**

497 D. Contini, M.R. Guascito, A. Mangone conceptualized the study design; A. Dinoi and E. Merico  
498 collected the samples and performed analysis of carbon content; A. Mangone and L.C. Giannossa  
499 performed chemical analysis. D. Cesari and D. Contini did statistical analysis and prepared the draft.  
500 All authors collaborated to interpretation of results, wrote, read, commented, and approved the final  
501 manuscript.

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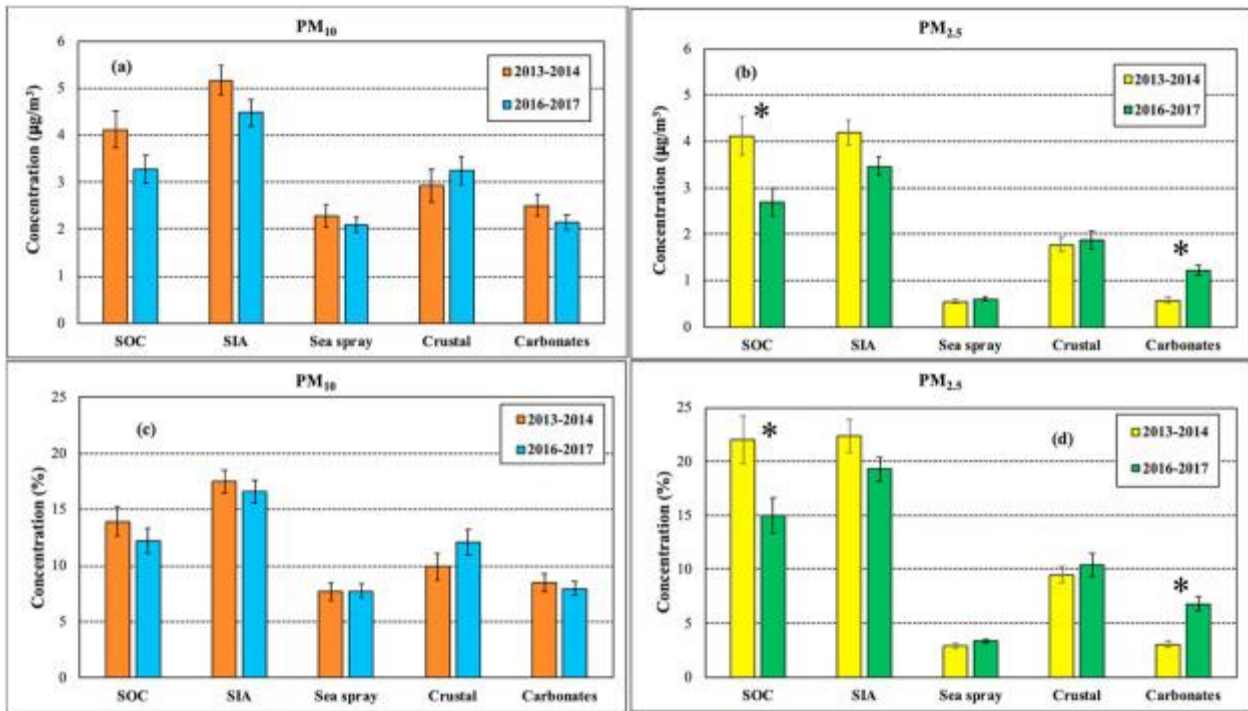


Fig. 1. Variability of absolute ( $\mu\text{g}/\text{m}^3$ ) and relative (%) concentrations of secondary species, SOC and SIA, sea spray, crustal and carbonates contributions in  $\text{PM}_{10}$  (a, c) and in  $\text{PM}_{2.5}$  (b, d). Errors bars represent the variabilities in terms of standard errors. \* p-value < 0.05 (ANOVA test).

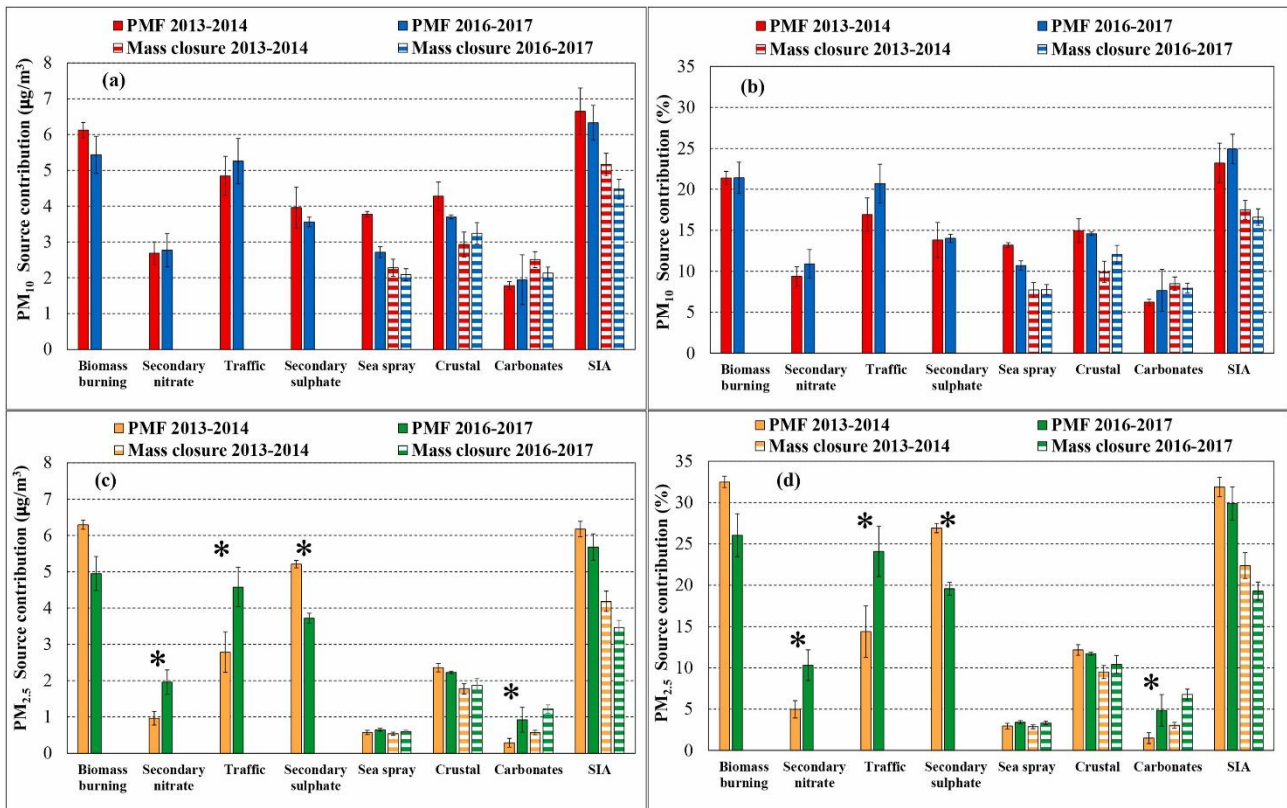


Fig. 2. Comparison of PM source contributions, in absolute (a, c), and relative terms (b, d), obtained with the two dataset 2013–2014 and 2016–2017. \* p-value < 0.05 (ANOVA test).

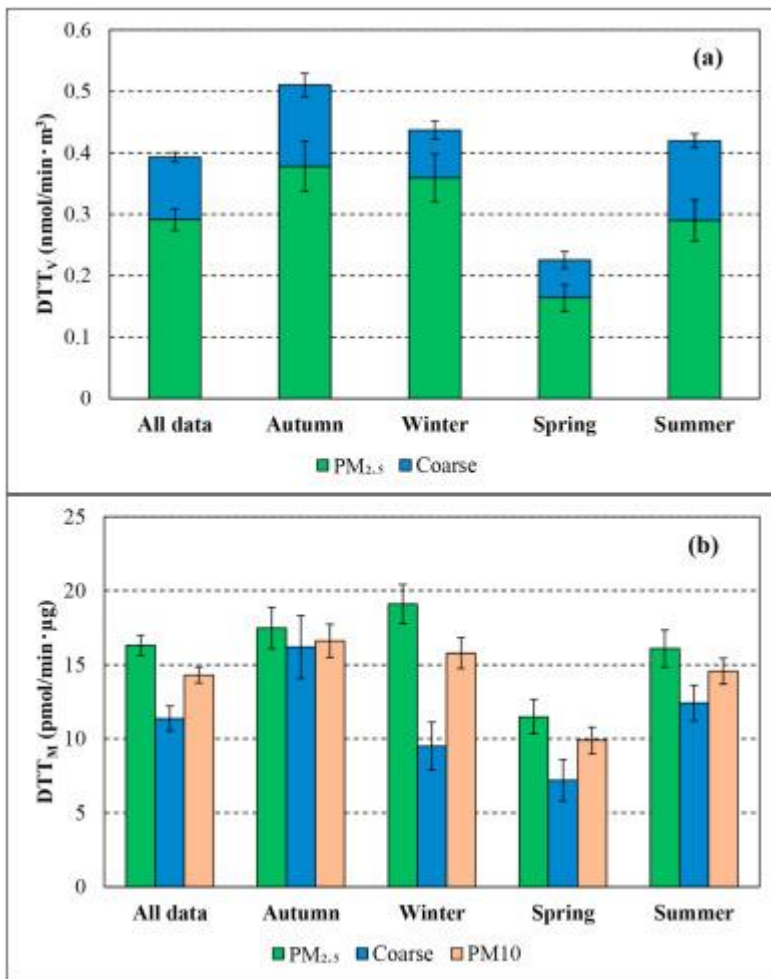


Fig. 3. (a) Seasonal variability of  $DTT_v$  ( $\text{nmol}/\text{min} \cdot \text{m}^3$ ) for  $\text{PM}_{2.5}$  and for the coarse ( $\text{PM}_{10-2.5}$ ) fraction. (b)  $DTT_M$  activity ( $\text{pmol}/\text{min} \cdot \mu\text{g}$ ) for  $\text{PM}_{2.5}$ ,  $\text{PM}_{10}$ , and coarse fraction for the entire 2016–2017 period. Errors bars represent the variabilities in terms of standard errors.



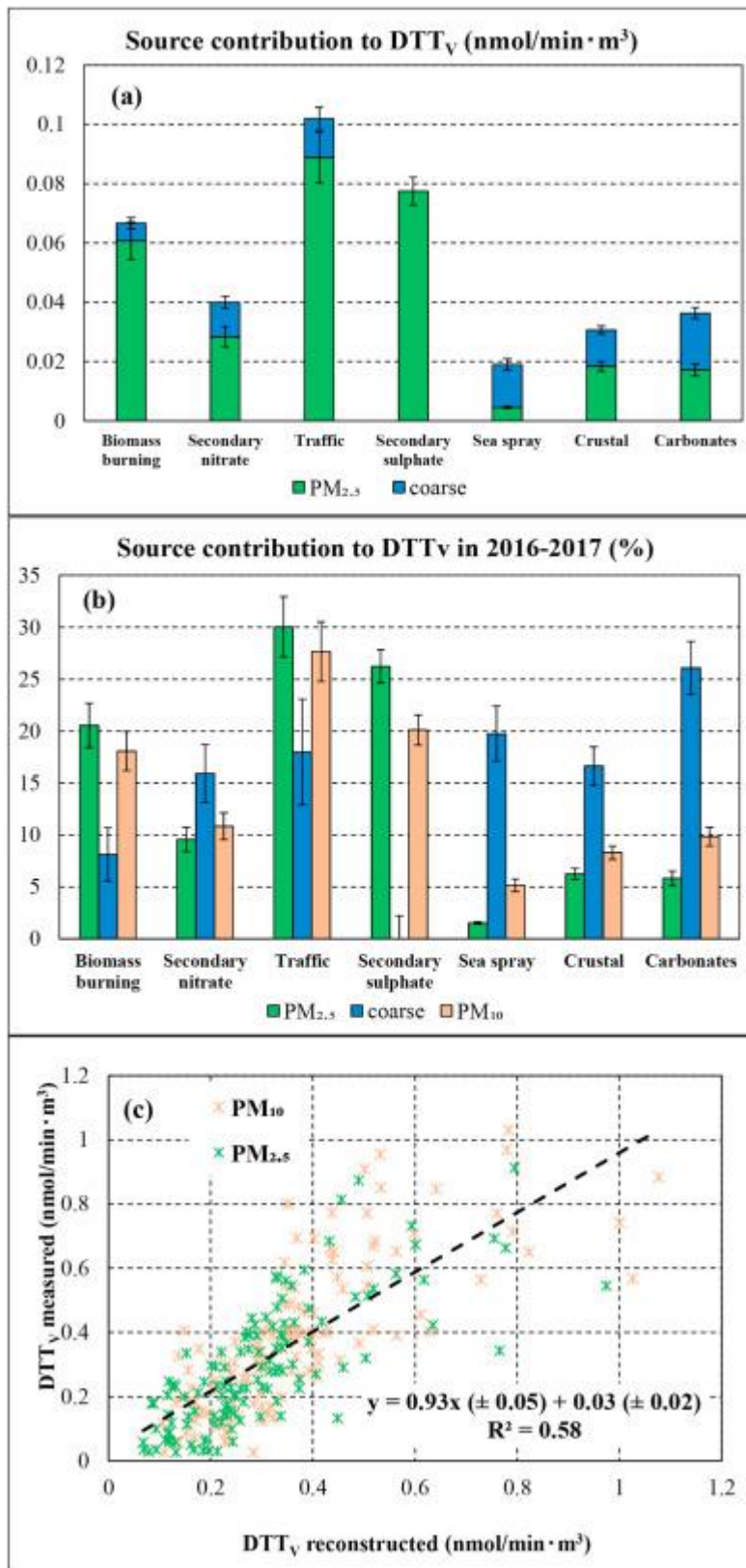


Fig. 4. (a) Absolute source contributions to DTT<sub>v</sub> in PM<sub>2.5</sub> and coarse fraction obtained using the MLR approach. (b) Source contributions in relative terms for PM<sub>2.5</sub>, PM<sub>10</sub>, and coarse fraction. (c) Comparison between measured and reconstructed DTT<sub>v</sub> including a linear fit. Errors bars represent the variabilities in terms of standard errors.