1	High temperature behavior of fedorite from Murun Alkaline Complex (Russia)
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28 Abstract

- 29 The thermal behavior of fedorite from the Murun massif (Russia) has been investigated by 30 means of differential thermal analysis (DTA), thermogravimetry (TG), in situ High Temperature 31 Single Crystal X-ray Diffraction (HT SCXRD) and ex situ High Temperature Fourier Transform 32 Infrared Spectroscopy (HT FTIR). Petrographic and Electron Probe Microanalyses (EPMA) 33 were also carried out. In thin section, fedorite (92 vol.%) is associated to quartz (5 vol.%) and 34 other minor phases (each < 1 vol.%). The morphogenetic features of minerals and their 35 paragenetic associations in the thin section allowed us to determine the sequence of mineral 36 crystallization from early to late ones: aegirine, apatite, amphibole, fedorite, microcline, rutile, 37 quartz, pectolite, and calcite. The crystal chemical formula of fedorite is: 38 $(Na_{1.56}K_{0.72}Sr_{0.12})_{\Sigma=2.40}(Ca_{4.42}Na_{2.54}Mn_{0.02}Fe_{0.01}Mg_{0.01})_{\Sigma=7.00}(Si_{15.98}Al_{0.02})_{\Sigma=16.00}(F_{1.92}Cl_{0.09})_{\Sigma=2.01}(O_{3.00})_$ 39 $7.93OH_{0.07}$ $\Sigma = 38.00 \cdot 3.2H_2O$. The TG curve provides a total mass decrease of ~5.5%, associated to the release of H₂O, H and F from 25 to 1050 °C. Fedorite crystallizes in the $P\overline{1}$ space group and 40 has: a = 9.6458(2), b = 9.6521(2), c = 12.6202(4) Å, $\alpha = 102.458(2), \beta = 96.225(1), \gamma = 12.6202(4)$ 41 119.902(1)° and cell volume, V = 961.69(5) Å³. The HT-SCXRD was carried out in air in the 42 43 25-600 °C T range. Overall, a continuous expansion of the unit cell volume was observed 44 although the c cell dimension slightly decreases from 25 to 600 °C. Structure refinements 45 indicated that the mineral undergoes a dehydration process with the loss of most of the interlayer H₂O molecules from 25 to 300 °C. The HT-FTIR spectra confirmed that fedorite progressively 46 47 dehydrates until 700 °C. 48 49 50 51
- 52 Keywords: fedorite, petrographic analysis, not-ambient, *in situ* HT-SCXRD, *ex situ* HT-FTIR.
 53

Ti; X = F, Cl, OH⁻, is a rare mineral occurring in alkaline rocks of Turiy and Murun complexes, 56 57 Russia (Kukharenko et al., 1965; Konev et al., 1993, 1996). This mineral was included in the 58 group of silicates with stacked modules of sheets of tetrahedra and octahedra (Ferraris, 1997). 59 Hawthorne et al. (2019) described fedorite as a sheet silicate with a structural unit consisting of two directions of infinite polymerization of tetrahedra, and an interstitial complex consisting of 60 four Ca²⁺ and Na⁺-octahedral sites with coordinations O_6^{2-} , $O_5^{2-}F^-$ (×2) and $O_4^{2-}F_2^{-}$, as well as 61 sites occupied by Na⁺ and K⁺ cations with positional disorder and coordinated by O²⁻ and partly 62 63 disordered H₂O molecules. 64 Difficulties in modelling interlayer cations and H₂O molecules were experienced in the structural studies of fedorite (Sokolova et al., 1983; Joswig et al., 1988; Mitchell and Burns, 2001; Kaneva 65 66 et al., 2020b). Elongated displacement ellipsoids were, initially, reported for one of the two interlayer cation sites (Joswig et al., 1988). The same sites, affected by strong anisotropy, were 67 68 split into A1 and A1a, A2 and A2a in structure refinement of the Turiy and Murun fedorite 69 (Mitchell and Burns, 2001). The refinement also showed: correlations between site occupancy factor and displacement parameters for one of the three H₂O sites, labelled *W*3; short distance 70 71 between W3 and A1a sites; short distance between A1a and A2 sites. Different hypotheses were 72 reported: W3 locally vacant when A1a is occupied; A2 locally vacant when A1a is occupied; A2 73 containing H₂O if locally occupied (Mitchell and Burns, 2001). Elongated displacement 74 parameters for interlayer cations and oxygen of the H₂O molecules were associated to mixed 75 cationic-anionic site occupancy in fedorite from three districts of the Murun complex (Kaneva et 76 al., 2020b). In particular, four K-occupied interlayer sites (K1 to K4) were identified: one (K2) or 77 more of them (K2 to K4) were split in two or three new positions. From one (O20w) to four 78 (O20w to O24w) H₂O molecules were also identified in the different analysed samples.

Fedorite, $A_{2-3}M_7T_{16}O_{38}X_2 \cdot nH_2O$, where $A = Na, K, Ba, Ca; M = Ca, Na, Mn^{2+}, Fe^{2+}; T = Si, Al,$ 55

79 The hydration degree of the alkaline cations in the interlayer of sheet silicate phases may

80 strongly affect the chemical and physical properties of these materials, especially in relation to

81 temperature changes. The dehydration process of fedorite was never explored so far.

82 In the present study, data from in situ High Temperature Single Crystal X-Ray Diffraction (HT-

83 SCXRD) as well as ex situ High Temperature Fourier Transform Infrared Spectroscopy (HT-

84 FTIR) experiments of fedorite are provided. Results from optical microscopy observations,

85 Electron Probe Microanalysis (EPMA), Differential Thermal Analysis (DTA),

86 Thermogravimetry (TG) and Derivative Thermogravimetry (DTG) are also reported.

87 This work is part of an ongoing research activity pointing to enhance the knowledge on the

stability of rare Russian alkaline minerals under non ambient conditions (Schingaro *et al.*, 2018;

- 89 Comboni et al., 2019; Lacalamita et al., 2019).
- 90

91 Experimental

- 92 Sample description and petrographic analysis
- 93 Fedorite-bearing sample Gav-33 was taken from brookite-quartz-feldspar rocks of the

94 Gavrilovskaya zone (same locality as sample Gav-43 from Kaneva et al., 2020b). The zone is

95 situated at the watershed of the Davan and Atbastakh Rivers, Irkutsk region, Bodaibo district

- 96 (58°19'51"N 119°5'45"E). The sample is an almost monomineral rock composed of scaly pale
- 97 pink fedorite (Fig. 1).

98 A thin section ~30 μm thick was studied by the optical petrographic method in transmitted light

99 using Olympus BX-51 polarized microscope. The morphogenetic features of fedorite and its

100 paragenetic associations were investigated, and the sequence of the crystallization was

101 determined from the relationships of minerals (Table 1).

102

103 Chemical analysis

104	A fedorite single crystal was embedded in epoxy resin, polished and carbon coated before
105	carrying out electron probe microanalysis. A JEOL JXA-8200 electron microprobe in full
106	wavelength dispersive spectrometry (WDS) mode was used. Operating conditions were: 15 kV
107	accelerating voltage, 5 nA sample current, ~ 1 μ m spot size and 40 s counting time. The used
108	standards were: grossular (Si, Al, Ca), omphacite (Na), olivine (Mg), K-feldspar (K), rhodonite
109	(Mn), fayalite (Fe), celestine (Sr), sanbornite (Ba), scapolite (Cl), horneblende (F).
110	A Phi-Rho-Z routine was employed for the conversion from X-ray counts to oxide weight
111	percentages (wt. %).
112	
113	Thermal analysis
114	Simultaneous differential thermal, thermogravimetric and derivative thermogravimetric analysis
115	(DTA/TG/DTG) were obtained by means of a Seiko SSC 5200 thermal analyzer equipped with
116	an ESS GeneSys Quadstar 422 quadrupole mass spectrometer which allowed the analysis of
117	gases evolved during thermal processes. The measurements were performed both in air at a

118 heating rate of 10 °C/min in the temperature range 25-1075 °C and in ultrapure helium at 20

119 °C/min in the heating range 25-1100 °C. Mass analyses were done in multiple ion

120 detection mode measuring the m/z ratios (ratio between the mass number and the charge of an

121 ion) 17 to 20 for H_2O and 19 for F.

122

123 X-ray diffraction analysis

124 Single crystals of fedorite were selected under an optical microscope and glued on the tip of a

125 glass fiber. The crystal (0.60x0.59x0.05 mm³) with the best diffraction behavior was used for the

126 in situ high temperature X-ray diffraction experiment by means of a Bruker AXS APEX II

127 diffractometer equipped with MoK α radiation ($\lambda = 0.71073$ Å) and a CCD area detector. The

- measurements were carried out in air from 25 to 600 °C by means of a Bruker AXS FR 559
- 129 home modified device (Zema *et al.*, 2022). A total of seven data collections were performed;

130 each data collection lasted ca. 11 hours. Before starting a new data collection, the crystal was 131 equilibrated for ca. 30 min. Data reduction was done using the software SAINT (Bruker, 2007); 132 empirical absorption corrections were applied by means of SADABS (Bruker, 2009); structure 133 refinements were performed by means of the program CRYSTALS (Betteridge et al., 2003) 134 using reflections with $I > 3\sigma(I)$ and starting from the atomic coordinates in Kaneva *et al.* (2020b). The structures were all solved in the triclinic space group $P\overline{1}$ (confirming the same 135 136 space group as found in Kaneva et al., 2020b). Overall scale factor, atomic positions, and atomic 137 displacement parameters were refined, starting from the coordinates of fedorite (Gav-43 sample) 138 in Kaneva et al. (2020b). At room temperature all cation and H2O molecules occupancies were 139 refined, whereas starting from 100 °C only the occupancies of the H₂O molecules were refined. 140 The occupancy of the tetrahedral site was constrained to 1; a restraint (1.00 ± 0.01) on the 141 occupancy of the Na plus Ca occupancies in the octahedral site was used to allow the occupancy 142 of this site to assume values higher or lower than 1. The attempts to anisotropically refine the 143 thermal ellipsoids of the interlayer Na cations and of the oxygen of the H₂O molecules result to 144 instable refinements. As a consequence, they were isotropically modeled. The analysis of the 145 difference-Fourier maps evidenced the presence of a residual electron density peak of about 1 $e^{-}/Å^{3}$ which may be considered as a consequence of the positional disorder of Na atoms. The 146 147 disorder modeling does not lead to a significant improvement of the refinement. Therefore, it 148 was not included in the final structure model. The effective coordination numbers (ECoN) of the 149 cations were calculated using the ECoN21 Software (Gheorghe, 2022) using a maximum 150 coordination radius of 4 Å. 151 Cif files of the refined crystal structures are submitted as supplementary material together with 152 details on structure analysis (Tables S1, S2).

153

154 Fourier Transform Infrared Spectroscopy

155 The infrared absorption spectra of fedorite were measured using an FT-801 spectrometer (Simex, 156 Novosibirsk, Russia). Powdered samples were mixed with anhydrous KBr, pelletized, and analyzed at a resolution of 1 cm⁻¹. A total of 32 scans were collected in the wavenumber range 157 158 500 to 4000 cm⁻¹. The IR spectrum of an analogous pellet of pure KBr was used as a reference. 159 IR variable temperature measurement, aimed at the study of fedorite dehydration were carried 160 out ex-situ, in the temperature range 25-700 °C, in steps of 50 °C following the procedure 161 described in Kaneva et al. (2020a) and Sapozhnikov et al. (2021) and here briefly summarized. 162 Fedorite and KBr were ground in a crucible, held at a given temperature for 5 min in a muffle 163 furnace, and then quenched to room temperature. Afterwards, two pressed pellets were prepared: 164 one consisting of a mixture of fedorite and KBr and another one of the pure KBr. Next, powders 165 of both samples were heated to the next temperature in the interval from 25–700 °C, and the 166 procedure was repeated. All pressed pellets were prepared and measured under the same 167 conditions of low relative humidity (<20%). In this way, water adsorption in heated samples can 168 be neglected.

169

170 **Results and Discussion**

171 *Petrography*

172 The host rock has a is holocrystalline coarse- and medium-grained structure and uniform texture

173 (massive). In the thin section, the sample mainly consists of fedorite (92 vol.%), with minor

174 components such as quartz (5 vol.%), microcline (less than 1 vol.%), aegirine (less than 1

175 vol.%), amphibole (less than 1 vol.%), calcite (less than 1 vol.%), pectolite (less than 1 vol.%);

accessory minerals are apatite and rutile (Fig. 2).

177 Fedorite is represented by elongated scaly subhedral grains up to $5.5 \times 1.2 \times 1.0$ mm in size,

178 showing perfect cleavage along {001}. Small euhedral crystals of aegirine $(0.07 \times 0.01 - 0.07 \times 0.01 + 0.07 \times 0.07$

179 0.05 mm), amphibole $(0.04 \times 0.02 - 0.10 \times 0.04 \text{ mm})$, and apatite $(0.02 \times 0.05 - 0.04 \times 0.09 \text{ mm})$

180 were found as inclusions in fedorite (Fig. 2c, d, k, l) which indicates its earlier crystallization

- 181 (Table 1). Fedorite grains are slightly cataclastic and can be replaced by secondary minerals
- 182 (probably, pectolite) along the cleavage and the cracks of the cataclase. Needle-like aegirine
- 183 grains are often located along the cleavage of the mineral (Fig. 2c). Larger subhedral aegirine
- grains $(0.4 \times 0.1 \text{ mm})$ are located at the boundary between fedorite and quartz grains (Fig. 2). In
- 185 thin section, aegirine grains exhibit pleochroism (from yellowish-green to blue-green).
- Larger subhedral grains of amphibole $(0.2 \times 0.15 0.07 \times 0.05 \text{ mm})$ and apatite $(0.4 \times 0.3 0.1 \times 0.05 \text{ mm})$
- 187 0.15 mm) also occur in xenomorphic relation to fedorite (Fig. 2g, j, l). In the thin section,
- amphibole grains are pleochroic from light green to gray, and the interference colors have an
- unusual ink color. Microcline grains up to 0.4×0.3 mm in size and anhedral quartz grains are
- 190 also xenomorphic in relation to fedorite (Fig. 2a, b, d-g, i).
- 191 In turn, quartz is xenomorphic with respect to the microcline. The grain sizes of quartz vary from
- 192 0.1×0.15 to 0.5×0.2 mm. Quartz contains inclusions of apatite and fedorite crystals up to 0.1×0.15
- 193 0.25 mm and 0.15×0.1 mm in size, respectively. At the grain boundary of quartz, single

elongated small grains of rutile up to 0.02×0.05 mm in size were found.

- 195 Calcite is rare. Its grains are quite small and do not exceed 0.05×0.05 mm. They are
- 196 xenomorphic in relation to fedorite.
- Fibrous grains of pectolite form fan-shaped aggregates and, possibly, were formed as a result ofhypergene processes.
- 199
- 200 Chemical composition
- 201 The average chemical composition of fedorite (wt%), determined over ten spots, is: SiO₂ (65.74),
- 202 Al₂O₃ (0.07), CaO (16.74), SrO (0.83), MnO (0.11), MgO (0.02), FeO (0.03), BaO (0.01), Na₂O
- 203 (8.58), K₂O (2.28), F (2.46), Cl (0.21). The crystal chemical formula, calculated on the basis of
- 204 16(Si+Al) and assuming the H₂O content derived by the structure refinement (see below), is:
- $205 \qquad (Na_{1.56}K_{0.72}Sr_{0.12})_{\Sigma=2.40}(Ca_{4.42}Na_{2.54}Mn_{0.02}Fe_{0.01}Mg_{0.01})_{\Sigma=7.00}(Si_{15.98}Al_{0.02})_{\Sigma=16.00}$

- 206 $(F_{1,92}Cl_{0,09}) = 2.01 (O_{37,93}OH_{0,07}) = 38.00 \cdot 3.2H_2O$. In the literature it is found that the number of H₂O 207 molecules may vary from 0.6 gpfu (sample Irk-53 in Kaneva et al., 2020b) to 3.69 gpfu (Turiy 208 fedotite in Mitchell and Burns, 2001). In addition, the Na/K ratio of our sample is close to that of 209 fedorite from Murun in Mitchell and Burns (2001) whereas fedorite in Kaneva et al. (2020b) 210 shows an enrichment in K atoms. 211 212 *TG/DTA/DTG data* 213 The TG-DTA-DTG experiments in air and He provided similar results (compare Fig. 3 and Fig. 214 S1). The DTA curve showed an exothermic peak at about 950 °C which may be associated to the 215 breakdown of fedorite. The thermogravimetric analysis, instead, exhibits a weight mass decrease
- of about 4% (650-700 $^{\circ}$ C T range) and 1.5% (1000-1050 $^{\circ}$ C). The former mass decrease is
- 217 compatible with the theoretical one associated to the loss of 3.4 H₂O gpfu. This value of the H₂O
- 218 content is close to that (3.2 H₂O gpfu) derived from the X-ray data (see below). The weight loss
- above 900 °C, instead, may be ascribed to the loss of fluorine as indicated by the measurement in
- the He atmosphere (Fig. S1).
- 221
- 222 *Crystal structure at room temperature*

223 Room temperature refinement of Gav-33 sample of fedorite was carried out in $P\overline{1}$ space group

and converged to $R_1 = 3.29$ %, $wR_2 = 4.01$ % (Table 2). Unit cell parameters and cell volume (*a*

225 = 9.6458(2), b = 9.6521(2), c = 12.6206(4) Å, $\alpha = 102.458(2)$, $\beta = 96.2250(10)$, $\gamma = 119.902(1)^{\circ}$,

226 $V = 961.69(5) \text{ Å}^3$, Table 2) are almost identical to those reported in literature (Mitchell and

227 Burns, 2001; Kaneva et al. 2020b). Atomic coordinates, site occupancy, and mean bond

distances at 25 °C are reported in Tables 3 and 4, respectively, in comparison with those

229 collected at 400 °C. A sketch of the mineral structure is, instead, reported in Fig. 4a, b. It is

- 230 characterized by eight symmetry independent tetrahedra (T1 to T8) which form regular
- 231 hexagonal rings of downward-pointing tetrahedra and distorted hexagonal rings of four

- 232 downward-pointing tetrahedra and two upward-pointing tetrahedra (T1 and T6) extending in the
- 233 *ab* plane (Fig. 4b). The *T*1 and *T*6 tetrahedra which share the apical oxygen atom, O19, show
- slightly short mean tetrahedral distances due to the low values of the Si-O19 individual distances
- 235 (Fig. 4a; Table 4). The structure of fedorite also consists of four symmetry independent
- 236 octahedra (M1 to M4) with mixed Na, Ca occupancy. Specifically, X-ray site-scattering
- refinement indicated a preferential partitioning of Na for the octahedral M1 site, an almost equal
- distribution of Na and Ca in M2, and the prevalence of Ca in M3 and M4 sites (Table 3). Similar
- 239 octahedral site occupancies were reported for other fedorite from Murun by Kaneva *et al.*
- 240 (2020b). The *M*1 octahedron which has $O_4^{2-}F_2^{-}$ coordination, exhibits the longest mean
- 241 octahedral bond distance and the highest bond length distortion (BLD_{M2} parameter) whereas the
- 242 *M*3 polyhedron, which is not coordinated by F atoms, is more regular (Table 4).
- 243 The interlayer site of the studied fedorite is occupied by four cation sites (*Na5* to *Na8*) with the
- 244 *Na5* showing a mixed Na/K occupancy whereas the *Na6* to *Na8* sites only contain Na atoms.
- 245 These cations coordinate five H₂O molecules, labelled from O20w to O24w, Table 3 and Fig. 4b.
- 246 The refinement also showed short K1-O20w, Na6-O21w, Na7-O24w, Na8-O22w, Na8-O23w
- 247 distances (2.041(4), 1.42(3), 1.50(2), 0.445(8), 2.01(3) Å, respectively) indicating a mixed
- 248 cation/anion occupancy. Short distances between H2O molecules (1.41-1.87 Å) point out to a
- statistical distribution of the molecules.
- 250 Similar octahedral site occupancies but different interlayer composition were reported for other 251 fedorite from Murun by Kaneva *et al.* (2020b). In the literature samples, indeed, the interlayer
- site is K-occupied and the K atoms, which are also affected by positional disorder, coordinate
- 253 from one to five H₂O molecules.
- 254

255 Crystal structure modifications and vibrational features upon heating

- 256 The thermal expansion of the fedorite structure proceeds from room temperature to $600 \text{ }^\circ\text{C}$
- 257 without change in symmetry $(P\overline{1})$ and with a variation of the cell volume which encompasses

258 the expansion along the *a* and *b*-axis but also the decrease of the *c* cell dimension (Fig. 5; Table 259 2). With increasing temperature, the mineral progressively dehydrated up to 300 °C when 260 fedorite lost most of the H₂O molecules as testified by the trend of the occupancies of O20w to 261 O24w sites (Tables 3, S1). In particular, the occupancy of O22w which is very close to the Na8 262 interlayer cation at RT (Fig. 6a), indicates that the relevant H₂O molecule is lost at 100 °C. The 263 O21w, O23w and O24w molecules, which are located in the *ab* plane around the O20w, tend to 264 shift toward the latter site (Fig. 6a-e; Table 5) which indeed is the only one occupied at T > 400265 °C (Table 3, S1). Fig. 7 shows that the total mean atomic numbers of the H₂O sites progressively 266 decrease with the temperature whereas the electron density of the O20w increases at T > 300 °C 267 as a consequence of the H₂O migration from the O21w, O23w and O24w sites towards the 268 O20w. At 400 °C a reduction of the effective coordination number (ECoN parameter, from 11 to 269 8) of the Na cations coordinating the H₂O molecules was also observed (Table 4). 270 The dehydration process proceeds from RT up to 400 °C causing a distortion of the hexagonal 271 rings of downward-pointing tetrahedra as testified by the shortening of the O8-O17 and O7-O15 272 internal diagonal distances (Fig. 6a, e; Table 5). The cations in the Na5/K1 site shift along the c 273 axis direction by moving away from the O20w and toward the F atoms, with a consequent 274 remarkable lengthening of the Na/K-O20w distance which is balanced by the shortening of the 275 Na/K-F distance (Fig. S2; Tables 4, S2). On the contrary, the cations in the Na6, Na7 and Na8 276 interlayer sites seem to move substantially in the *ab* plane, being their *z* coordinates almost 277 unchanged during heating (Table 3, S1). In particular, they progressively increase their distance 278 from the O20w (Fig. S2). In addition, the cation in Na6 moves away from the O10 atom and 279 approximates the O14; the cation in Na7 moves away from the O16 and O17 and shifts toward 280 the O11 and O13; the cation in Na8 moves from the O7 and O15 atom toward the O16, O18 and 281 O19 oxygens (Fig. S2, 6f). This also results in the shortening of the distances between equivalent 282 Na cations (Fig. 6f).

283 Finally, structure refinements of the partially dehydrated fedorite (from 400 to 600 °C) evidence

that no significant geometric modifications of the tetrahedral and octahedral sheets were induced

285 by the heating process, since quite similar values were obtained for tetrahedral and octahedral

286 mean distances and geometrical parameters (Tables 4, S2).

287 The infrared spectra of the fedorite sample annealed at different temperatures (from 25 to 700

[°]C) are given in Fig. 8. The group of peaks observed at 1119, 1030, 790, 618 cm⁻¹ may be

assigned to the asymmetric and symmetric Si–O stretching modes of the SiO₄ tetrahedra.

290 The peak at 618 cm^{-1} in the studied sample is slightly shifted in comparison with the same peak

in Gav-43 and Irk-53 samples (615 cm⁻¹; Kaneva *et al.* 2020b). The energy of bending vibration

292 of SiO₄ at 615 cm⁻¹ is slightly increased in the studied sample which contains higher H₂O

293 concentration. The same behavior of SiO₄ bending vibrations has been observed in elpidite

samples in Bogdanov *et al.* (2021).

295 Intensities of H–O–H asymmetric and symmetric stretching vibrations in the region 3000–3700

296 cm⁻¹ (in the inset of Fig. 8) and bending ones at 1627 cm⁻¹ decreased during annealing. This

indicates that the fedorite sample was losing H_2O molecules. Also the peak at 790 cm⁻¹ in initial

sample shifts to 796 cm⁻¹ in annealed at 700 °C sample when the fedorite has lost almost all the

299 H₂O molecules. It could be due to slightly change in Si–O–Si angles during annealing (Shendrik

300 *et al.*, 2021).

301 The region $3000-3700 \text{ cm}^{-1}$ of absorption spectrum related to stretching vibrations of H₂O

molecules has a complicated shape (Fig. 8). Six peaks at 3177, 3429, 3549, 3588, 3635, and

303 3656 cm⁻¹ can be clearly seen in the sample at room temperature. Such peaks are similar to those

found in the Irk-53 sample (Kaneva *et al.* 2020b). Three groups of peaks can be distinguished on

305 the basis of the behavior during heating (Fig.s 8, 9). Group 1 includes the 3588 and 3656 cm^{-1}

306 peaks. The intensity of these peaks decreases as the temperature increases, becomes negligible at

temperatures above 327 °C (Fig. 9). Group 2 includes the 3177, 3429, 3549, and 3635 cm^{-1} .

308 Their intensity gradually decreases upon annealing and becomes negligible above 627 °C (Fig.

309 9). Group 3 includes 3576 and 3650 cm⁻¹ peaks. The latter start to be observed at above 150 $^{\circ}$ C 310 and show the maximum intensity at 350 °C (Fig. 9). Further annealing at higher temperatures 311 leads to a decrease in their intensity. Peaks belonging to Group 1 and 3 are sharp and may be attributed to OH-anions in different positions or to O-H stretching vibrations (hydrogen bonds of 312 313 medium strengths). Group 2 bands may be attributed to H₂O molecules with a strong hydrogen 314 bond. The trends of the band intensities (Fig. 9) in the FTIR spectra may be explained by taking 315 into account the results of the HT-SXRD structure refinements. Specifically, the quick 316 decrement of the intensity observed for the Group 1 agrees with the reduction of the total water 317 means atomic numbers until ~ 300 °C (Fig. 7). The intensity variation of the Group 3 seems to 318 describe the migration of the H₂O molecules at $T \ge 100$ °C (Fig. 6). Finally, the gradual 319 decreasing of the band intensity of Group 2 well represents the evolution of the dehydration 320 process affecting the fedorite up to T ~ 600 $^{\circ}$ C as described above (see the *Crystal structure* 321 *modifications upon heating* section).

322

323 Conclusions

324 The combination of thermal, in situ HT-SCXRD and ex situ HT-FTIR analyses allowed to detail 325 the thermal behavior, and specifically, the dehydration process of fedorite from Murun massif. 326 The mineral contains about 3 H₂O gpfu which occupy five different crystallographic sites (O20w 327 to O24w) in the structure refined at room temperature, while at 600 °C only the O20w resulted 328 occupied. Fedorite underwent thermal expansion in the whole temperature range (25-600 °C) 329 explored by XRD with a preferential dilatation along the *a* and *b* axes while a decrement of the *c* 330 cell parameter was a consequence of the dehydration process. All the used analytical techniques 331 testified for a continuous dehydration reaction up to T ~ 600 $^{\circ}$ C whereas the complete loss of 332 H₂O molecules at 650-700 °C was highlighted by the TG curve. The combination of the XRD 333 and FTIR results allowed to detail the features of the partial dehydrated structures. In particular, 334 one H₂O molecule (O22w) early left the interlayer site whereas the reduction in concentration of

the remaining molecules rapidly proceeded up to T ~ 300 °C. This process was accompanied by
the migration of the molecules in O21w to O24w toward the inversion center (O20w) of the
crystal structure. The dehydration also caused the decrease in the coordination of the interlayer
Na cations which yielded their shift toward their symmetry equivalents. The release of the F
atoms occurred above 700 °C (TG curve) whereas the breakdown of the structure at about 1000
°C was pointed out by the DTA curve.

341

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354

355 **Competing interests.** The authors declare none.

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408 FIGURE CAPTIONS

409 Fig. 1. Sample of fedorite from the Gavrilovskaya zone (sample Gav-33): (a) unpolished sample;
410 (b) polished sample.

- 411 Fig. 2. Thin section image of sample Gav-33 in transmitted light. (a, d, e, g, i, l) polarizers are
- 412 parallel; (b, c, f, h, j, k) polarizers are crossed. The interference color corresponds to a thickness
 413 of thin section (~ 30 μm).
- 414 Fig. 3. Differential thermal analysis (DTA, black line), thermogravimetric analysis (TG, blue
- 415 line), and the TG first derivative (DTG, red line) of Gav-33 sample measured in air.
- 416 **Fig. 4.** (a) Crystal structure of the fedorite sample, Gav-33; (b) a detail, as seen along [001], of
- 417 the tetrahedral (orange), octahedral (blue) and of the interlayer site which is occupied by Na⁺
- 418 (dark yellow) and K⁺ (light yellow) cations. Oxygen and fluorine atoms are represented in red
- 419 and green, respectively.
- 420 Fig. 5. Normalized unit cell parameters and volume of the studied fedorite *versus* temperature.
- 421 *a*₀, *b*₀, *c*₀ and V₀ are cell parameters and volume at room temperature, respectively. The size of
- 422 the symbols is larger than the associated esd's.
- 423 Fig. 6. (a-e) Interlayer cavity of fedorite from refined structure in the RT-600 °C T range; (f)
- 424 extended interlayer cavity view at 400 °C. Colors as reported in Fig. 4; gray arrows indicate the
- 425 shift of Na cations with the temperature.
- 426 **Fig. 7.** Mean atomic numbers (m.a.n.'s, e⁻) of the H₂O molecules *versus* temperature.
- 427 Fig. 8. IR absorption spectra of fedorite annealed at different temperatures. In the inset, the
- 428 region of O-H stretching vibrations is given.
- 429 Fig. 9. Temperature dependence as a function of the relative intensity of IR absorption of
- 430 different OH stretching vibrations of annealed fedorite sample.